

AOS 801: Advanced Tropical Meteorology  
*Lecture 10 Spring 2023*  
Vertical Velocity, Precipitation,  
Buoyancy and Water Vapor 2

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# Announcements

Paper discussion today, let's do it 25 minutes before class ends.

HW2 is due March 6.

We are finishing our discussion about tropical convection today.

# Leading Balance in Convection

We can use mass continuity to merge the two equations to obtain the following:

$$\frac{\partial^3}{\partial z^3} \left( \frac{1}{\rho} \frac{\partial \rho w}{\partial z} \right) \simeq - \frac{\rho}{\mu_c} \nabla_h^2 B$$

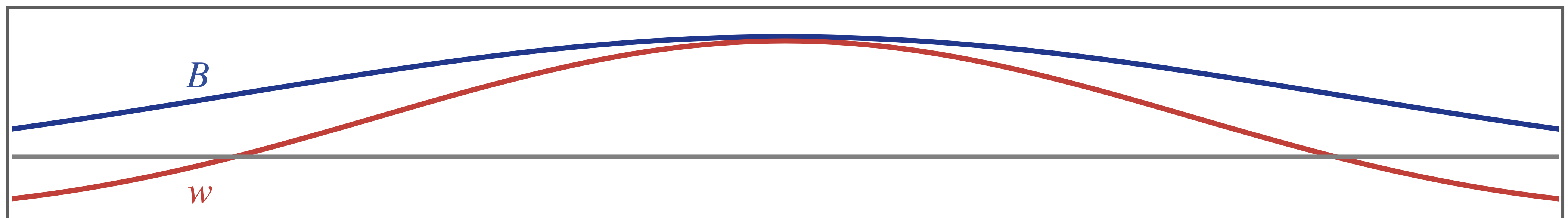
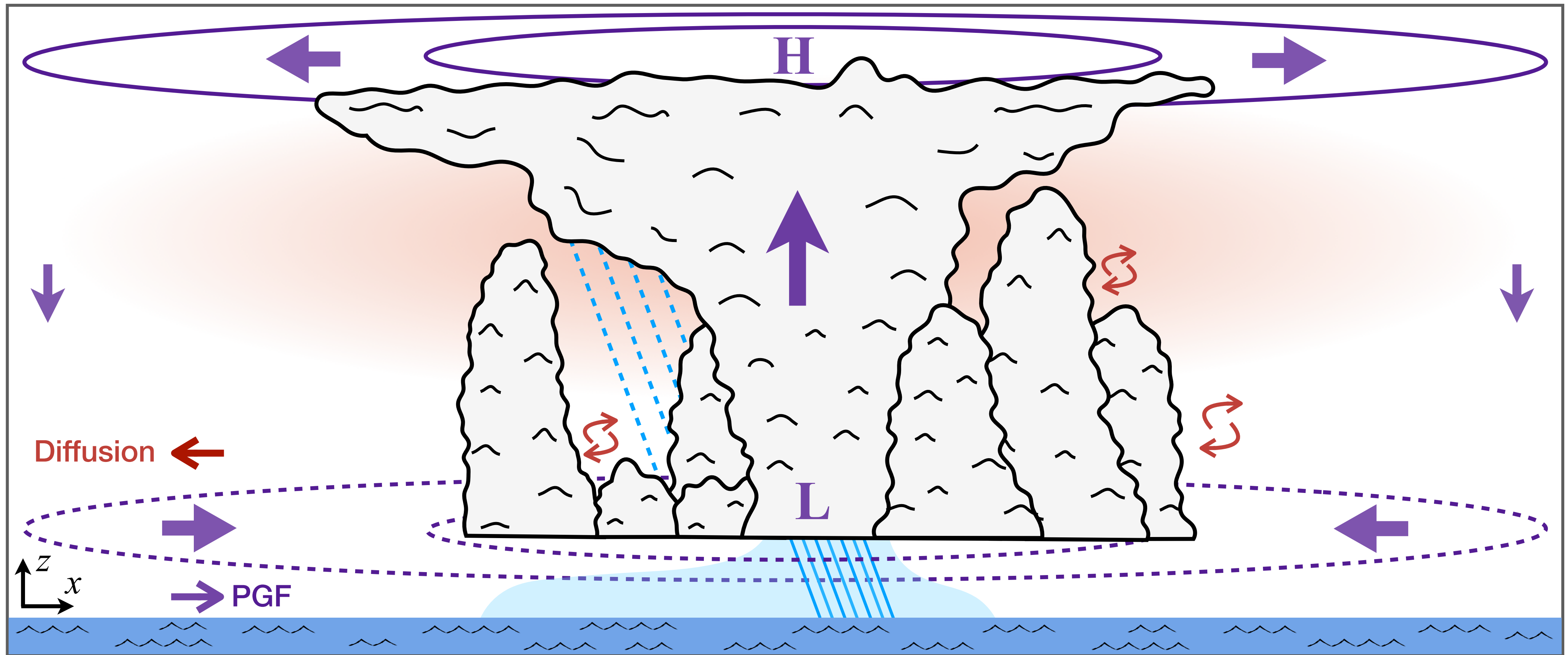
In pressure coordinates you get an even simpler relation:

$$\frac{\partial^4 \omega}{\partial p^4} = \frac{1}{\mu_c^*} \nabla_h^2 \alpha'$$

$$\alpha' = R_d T' / p \quad T' = T_c - T_0 \quad \mu_c^* = \omega_c p_\ell \sim g^2 \mu_c \sim 10^5$$

What does this mean?

# Sticky Convection



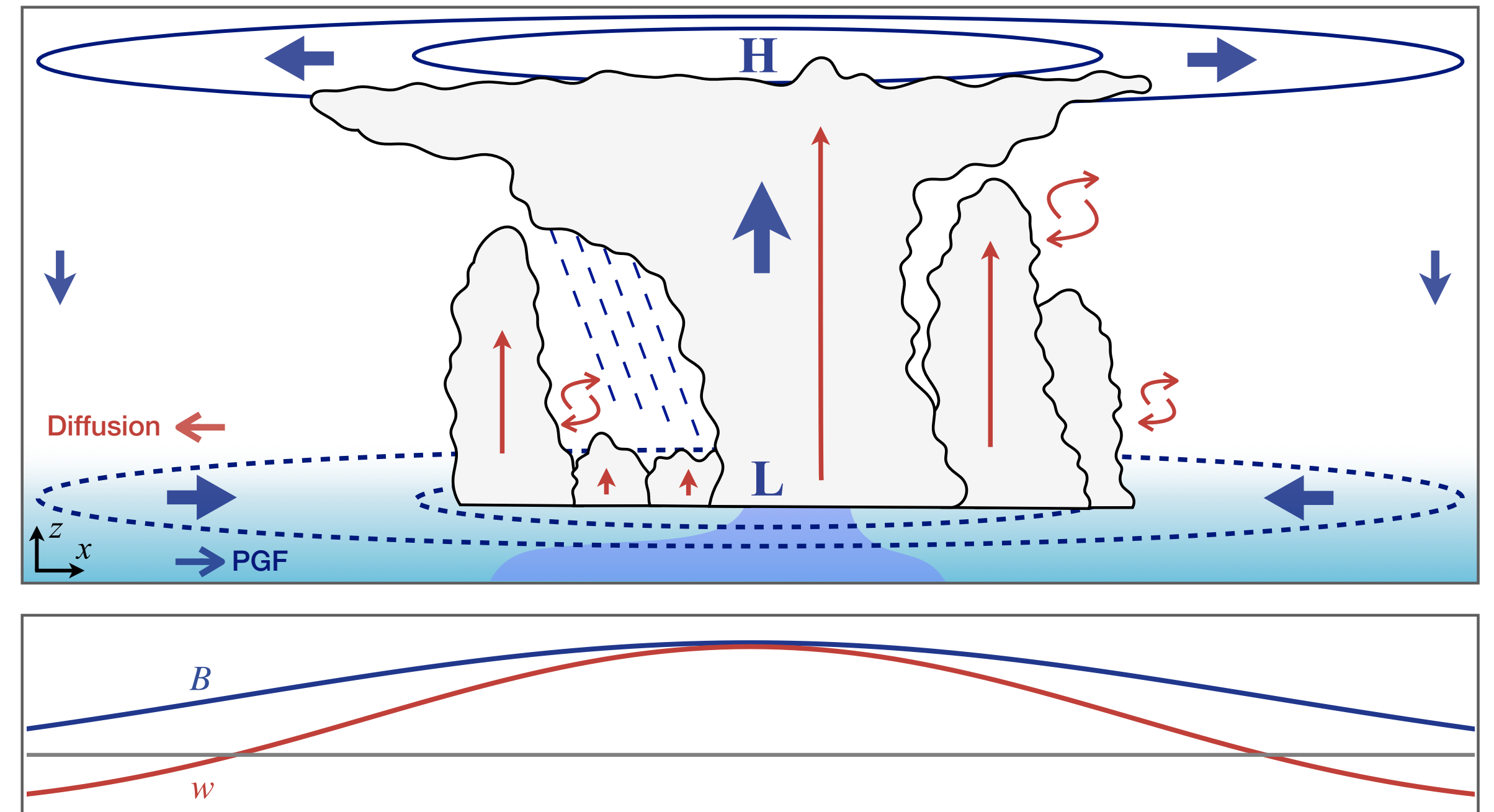
# Sticky Convection

Individual Cloud elements are turbulent but their aggregate behavior is not.

Antitriptic balance implies that larger-scale flow is quasi-laminar.

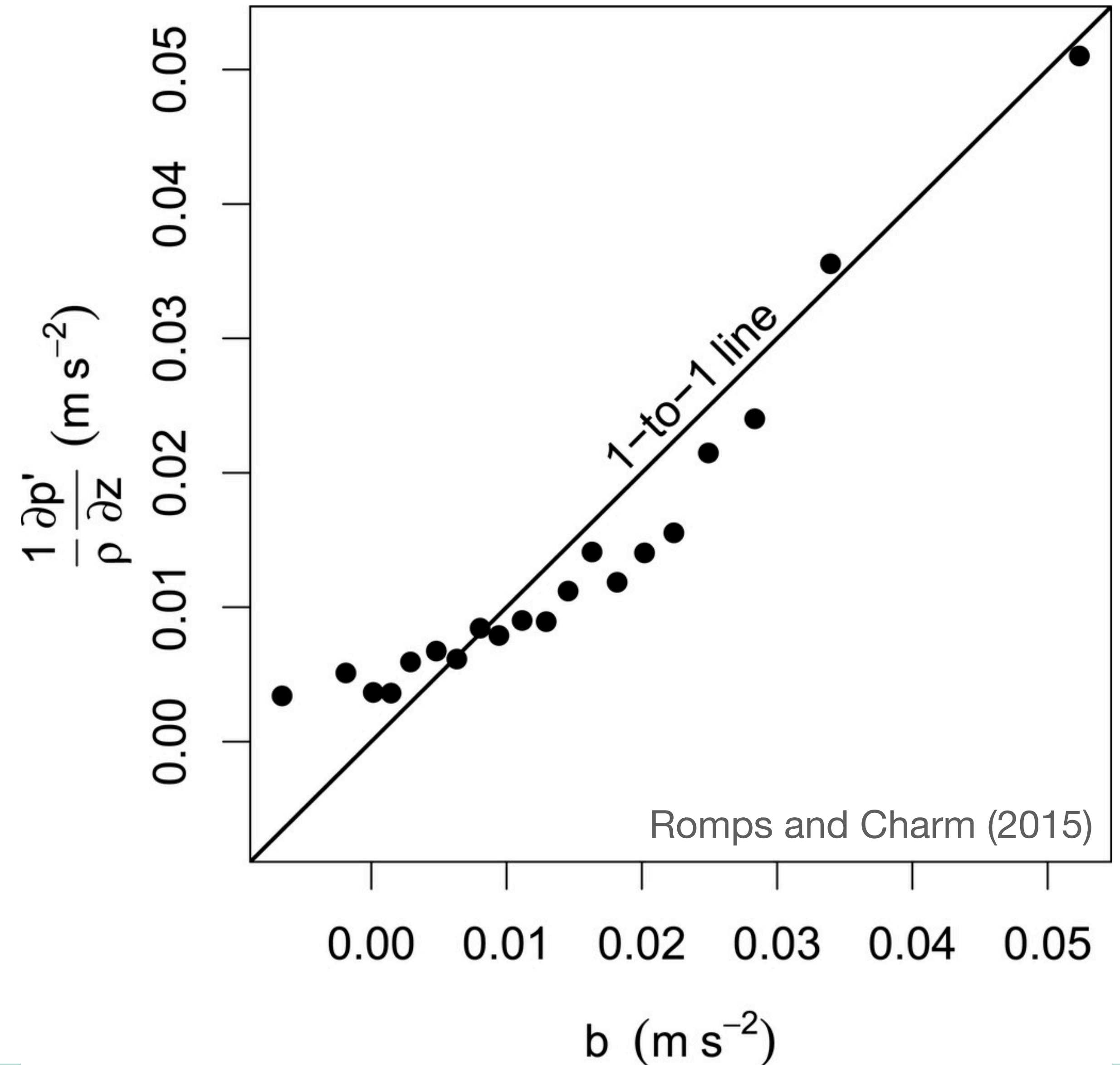
The winds flow down the pressure gradient and rise in the region of high buoyancy.

There is a direct relation between vertical velocity and buoyancy.



# Sticky Thermals

This relationship applies even for shallow cumulus clouds both in and out of the tropics



# Precipitation-buoyancy relation

Assuming rigid lids at the surface and tropopause, we can simplify this to

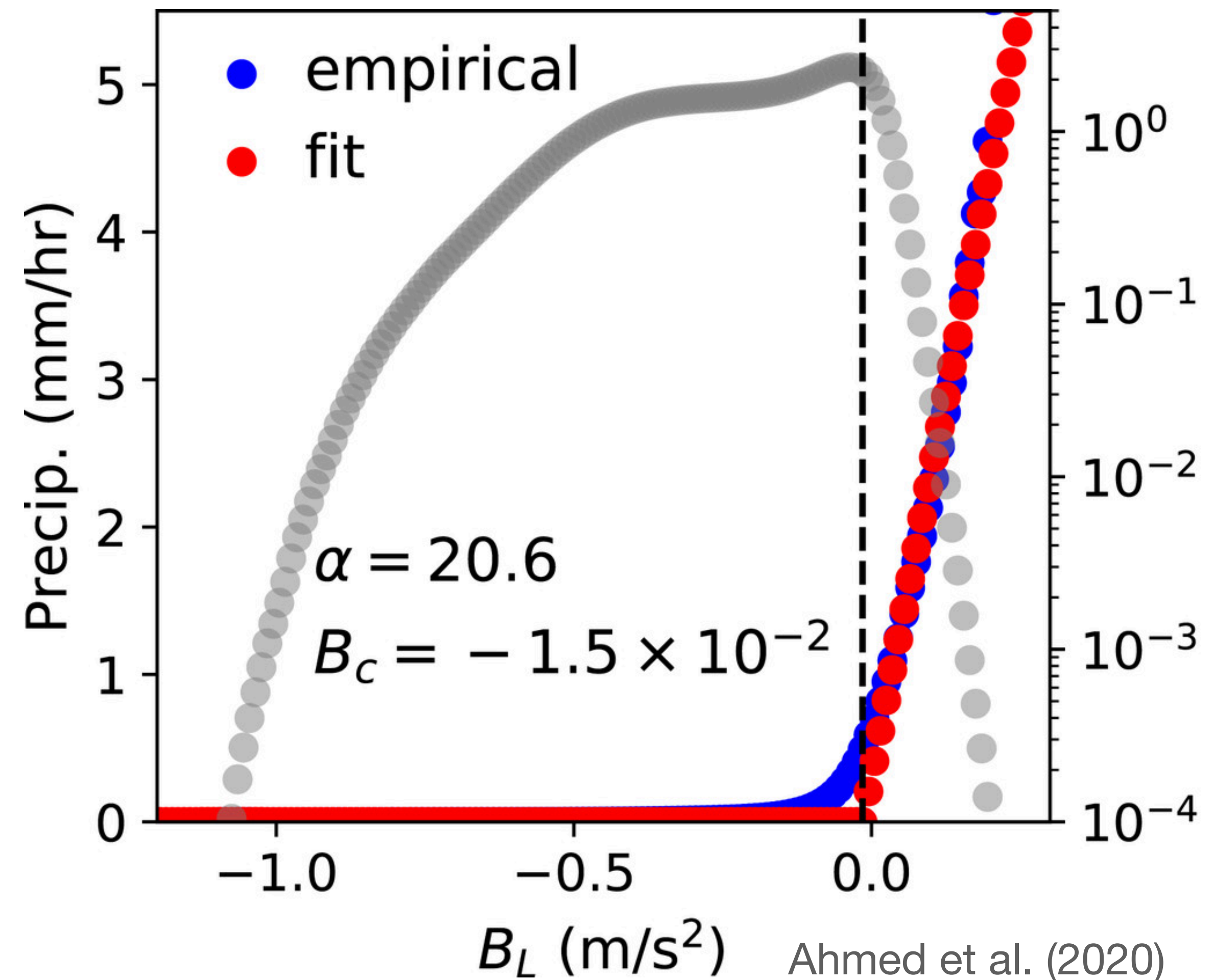
$$\mu_c^* m^4 \langle \omega \rangle = \nabla_h^2 \langle \alpha' \rangle$$

Where  $m$  is a vertical wavenumber. We now invoke the WTG approximation:

$$\langle \omega \rangle S_p = L_v P + \langle Q_r \rangle$$

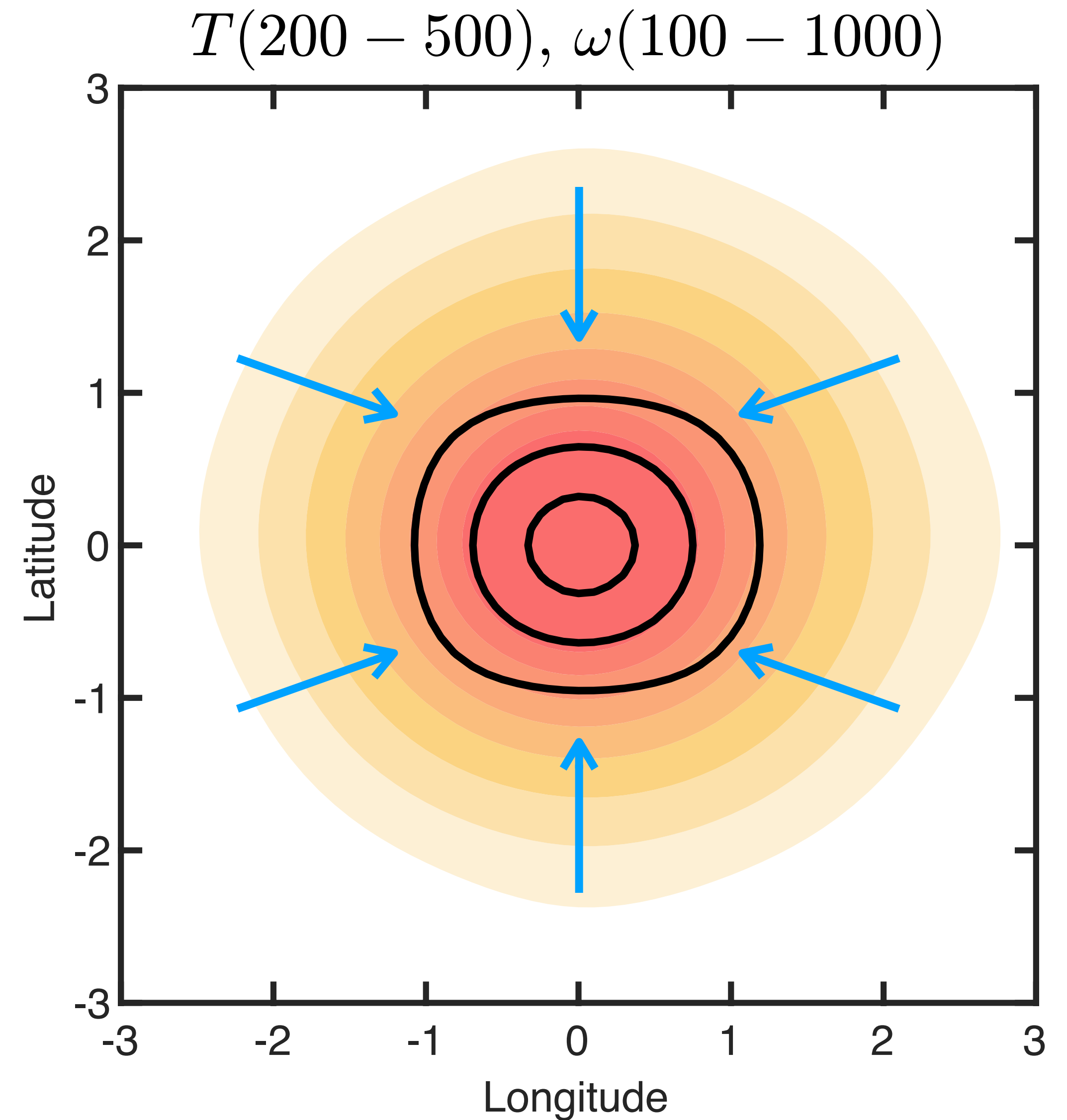
Combining everything we have:

$$L_v P = \frac{S_p}{\mu_c^* m^4} \nabla_h^2 \langle \alpha' \rangle - \langle Q_r \rangle$$



# ERA5 relation between $T$ and $w$

What if the individual convective elements are diffusive and entrain turbulently, but the entire region of ascent behaves like a plume that entrains dynamically ?

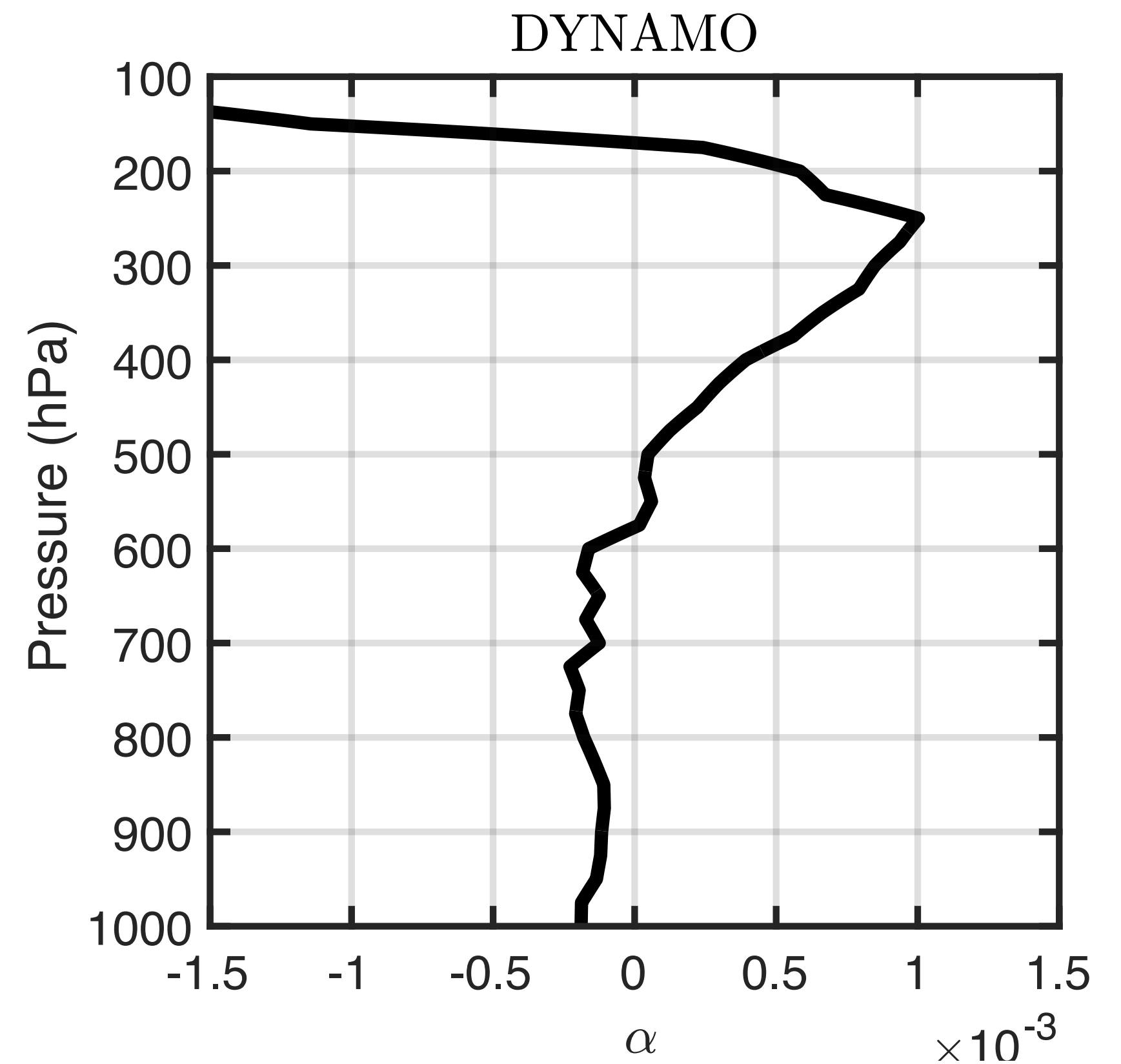
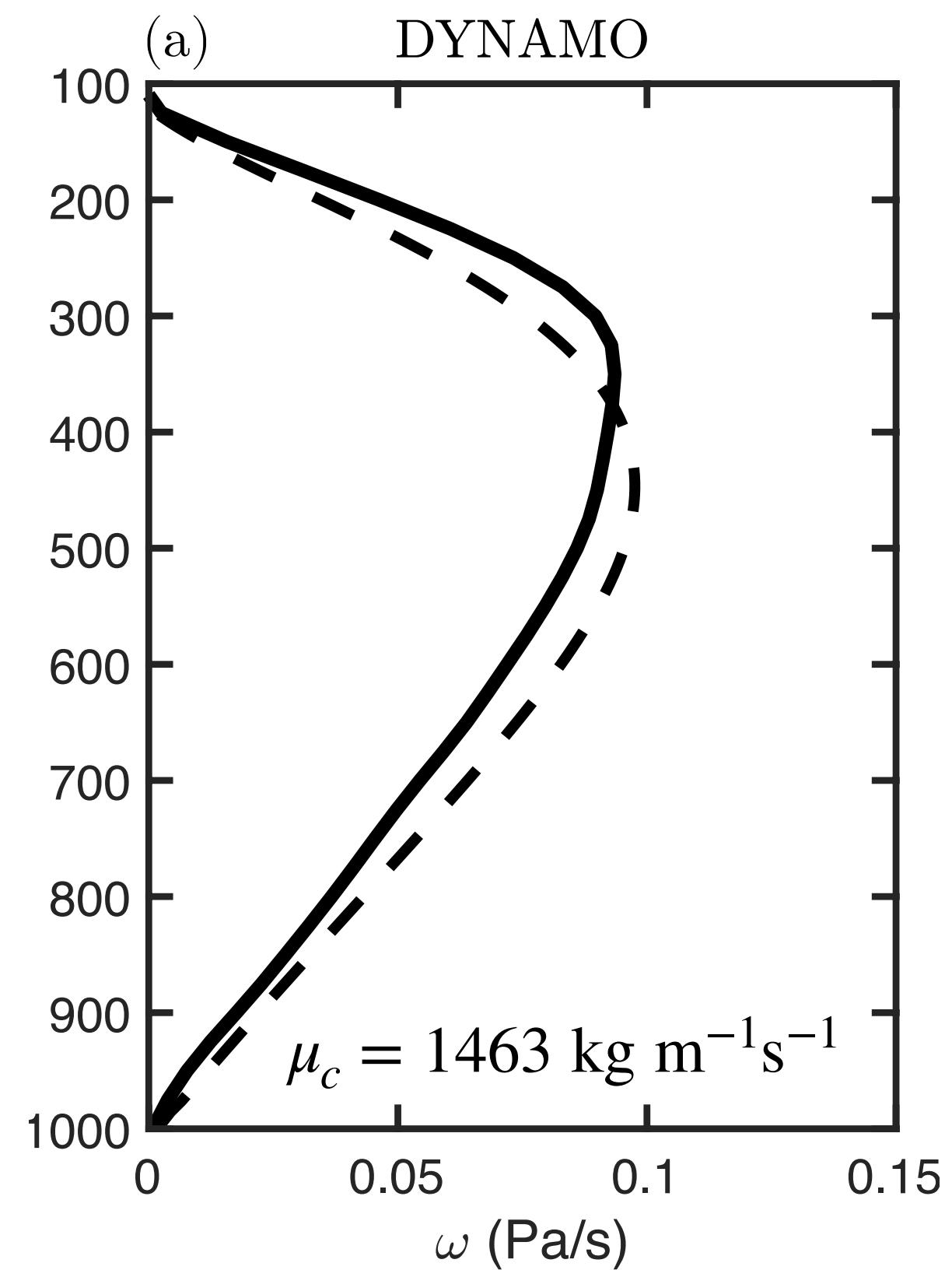




# Relationship between alpha and omega

What if we assume vertical wave solutions ( $m$  is a vertical wavenumber) with rigid lids at the surface and tropopause:

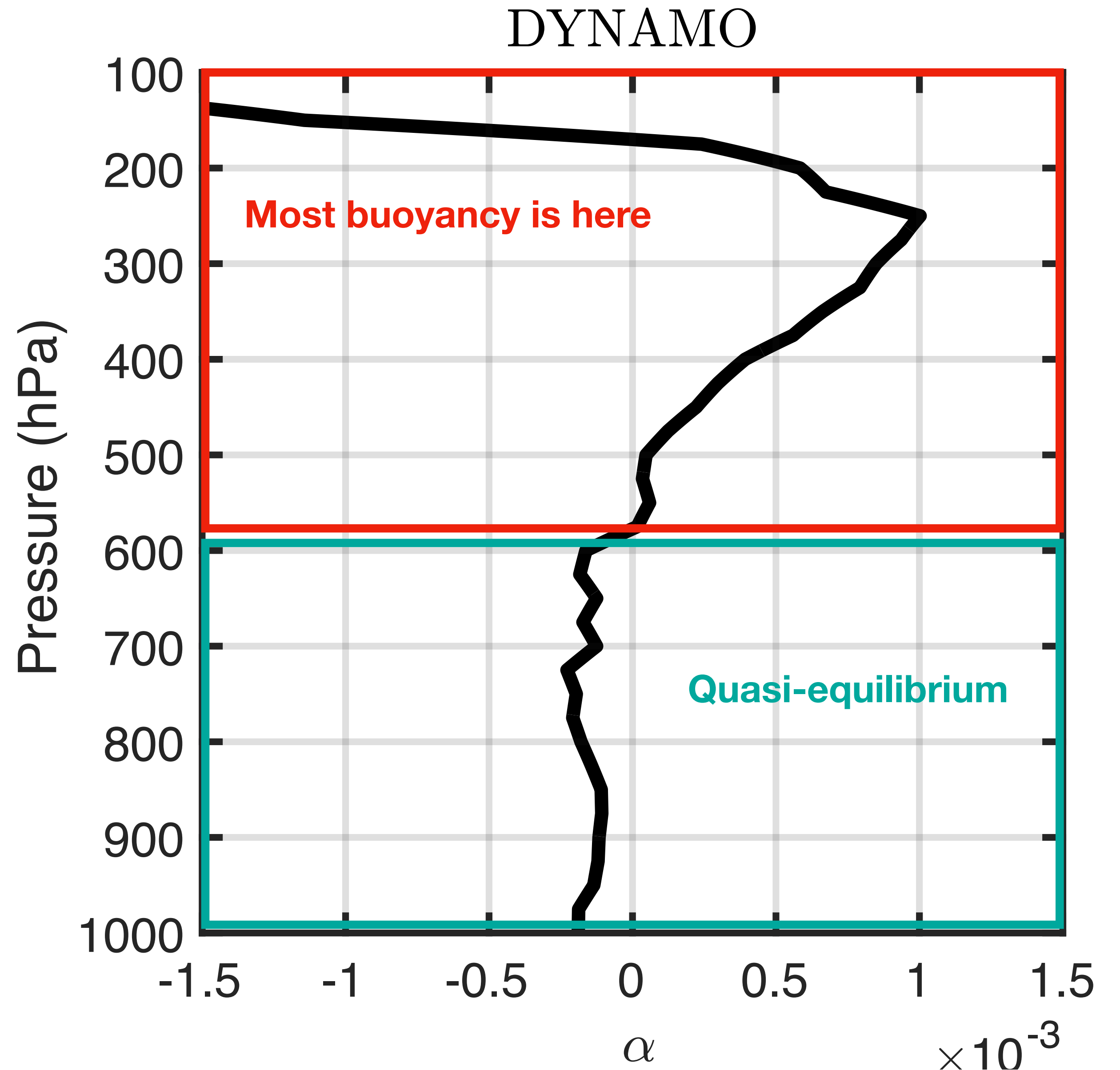
$$\mu_c^* m^4 \omega = \nabla_h^2 \alpha'$$



# Relationship between alpha and omega

$$\mu_c^* m^4 \omega = \nabla_h^2 \alpha'$$

Let's try dividing  $\alpha'$  into two regions: a lower troposphere with small values, and an upper troposphere with large values.



# Dynamic entrainment

The mesoscale plume entrains dynamically:

$$\frac{\partial \text{MSE}_c}{\partial z} = -\epsilon_D (\text{MSE}_c - \text{MSE}_e).$$

Assuming that the plume is saturated we have:

$$\text{MSE}_c(p_{fl})^* = -\frac{1}{\omega_c(p_{fl})} \int_{p_{fl}}^{p_s} \frac{\partial \omega_c}{\partial p} \text{MSE}_e dp.$$

If the mass flux is assumed linear beneath the freezing layer we have

$$\text{MSE}_c^*(p_{fl}) \simeq \text{MSE}_{eL}.$$

The MSE that makes it to the top of the troposphere is the average MSE beneath the freezing layer

# Special case

Assuming latent energy dominates changes in lower troposphere MSE, and a super small BL we get:

$$\langle \omega \rangle = \frac{L_v}{\mu_c^* m^4 \Delta p} \ln \left( \frac{T_{fl}}{T_{lnb}} \right) \nabla_h^2 \langle q \rangle$$

Vertical velocity is determined by the laplacian of column moisture. Using the WTG approximation we have

$$\langle \omega \rangle S_p = L_v P + \langle Q_r \rangle$$

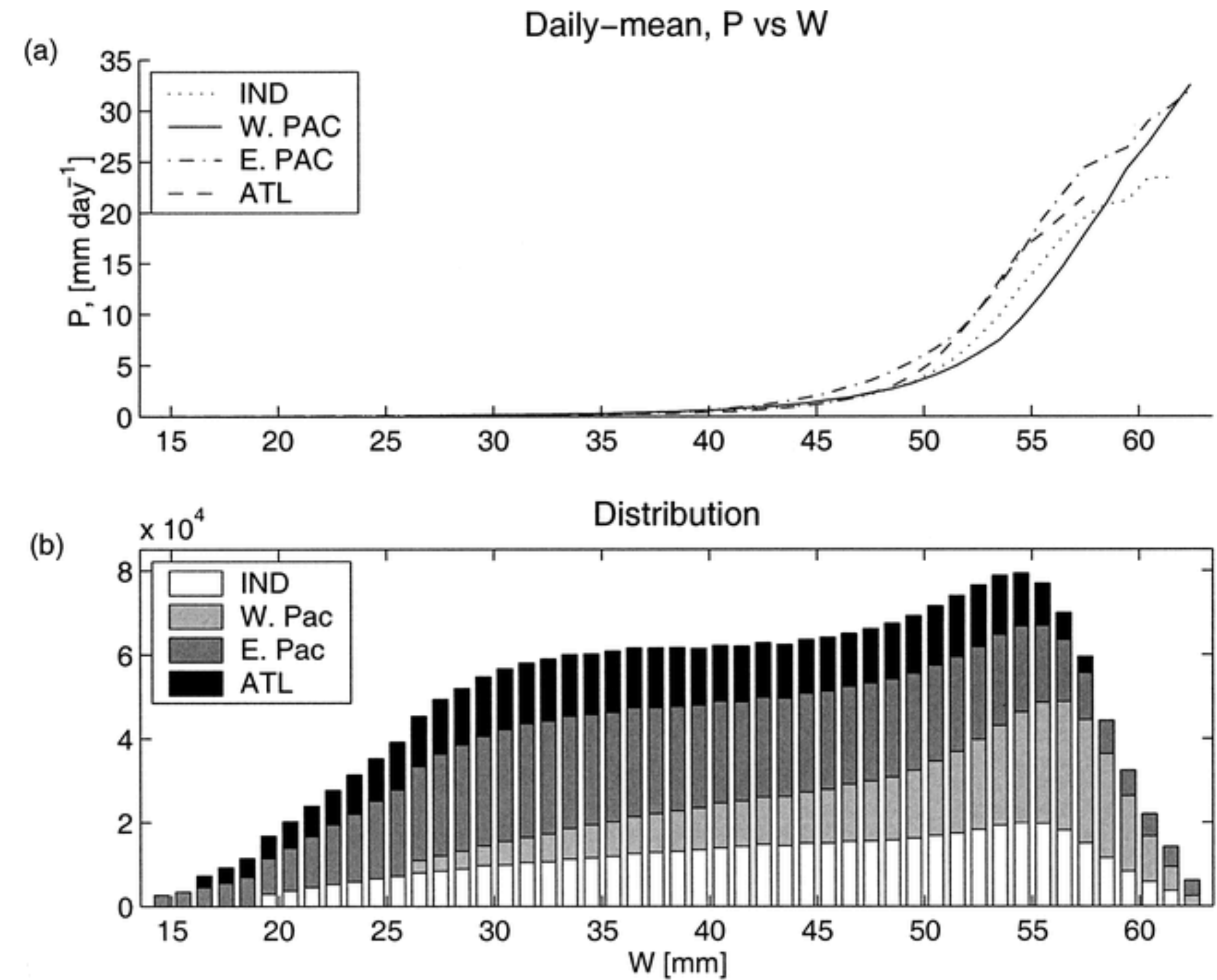
So we get

$$L_v P = \frac{S_p L_v}{\mu_c^* m^4 \Delta p} \ln \left( \frac{T_{fl}}{T_{lnb}} \right) \nabla_h^2 \langle q \rangle - \langle Q_r \rangle$$

# Precipitation-moisture relation

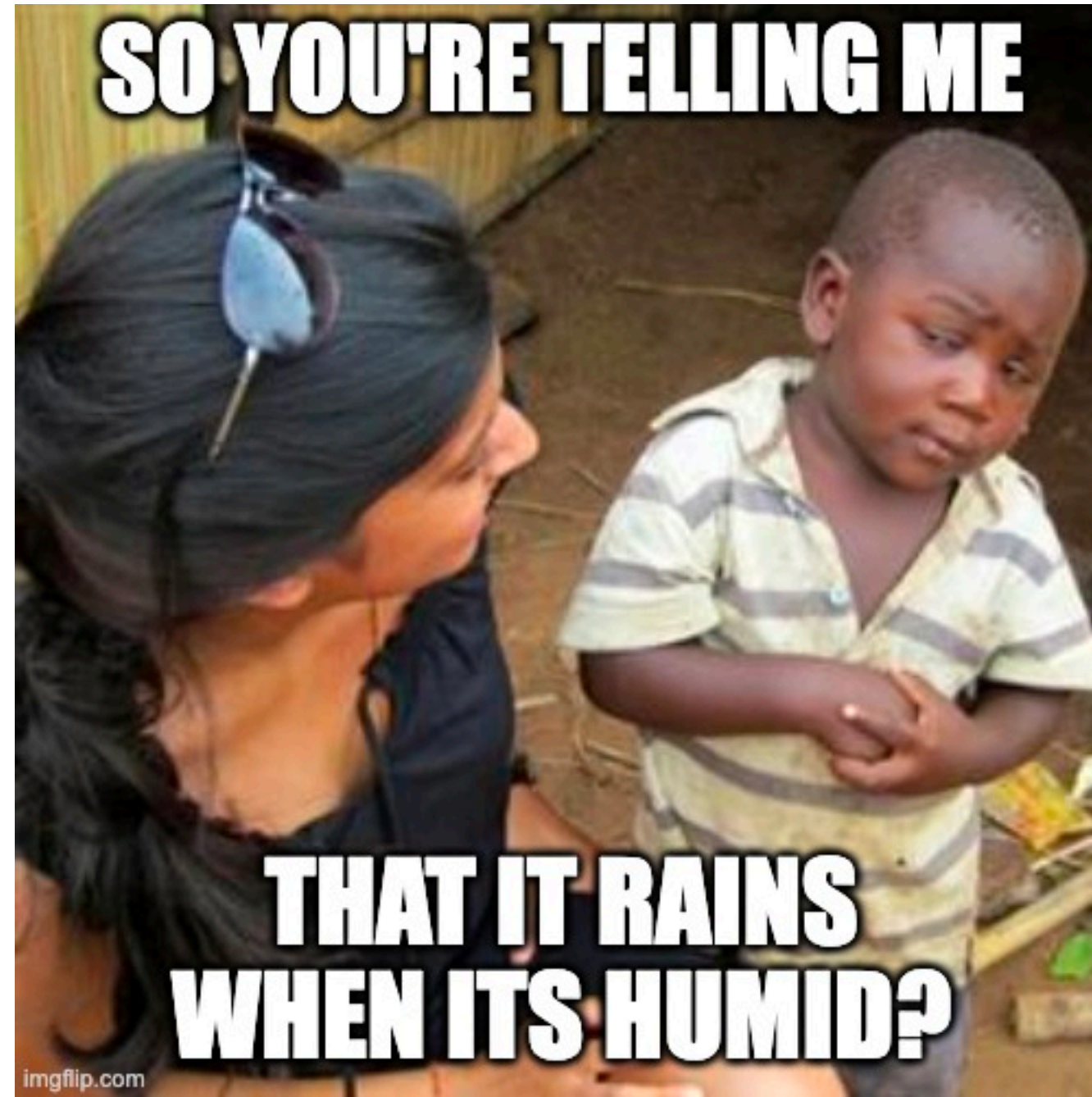
$$L_v P = \frac{S_p L_v}{\mu_c^* m^4 \Delta p} \ln \left( \frac{T_{fl}}{T_{lnb}} \right) \nabla_h^2 \langle q \rangle - \langle Q_r \rangle$$

Precipitation is a function of the Laplacian of column water vapor



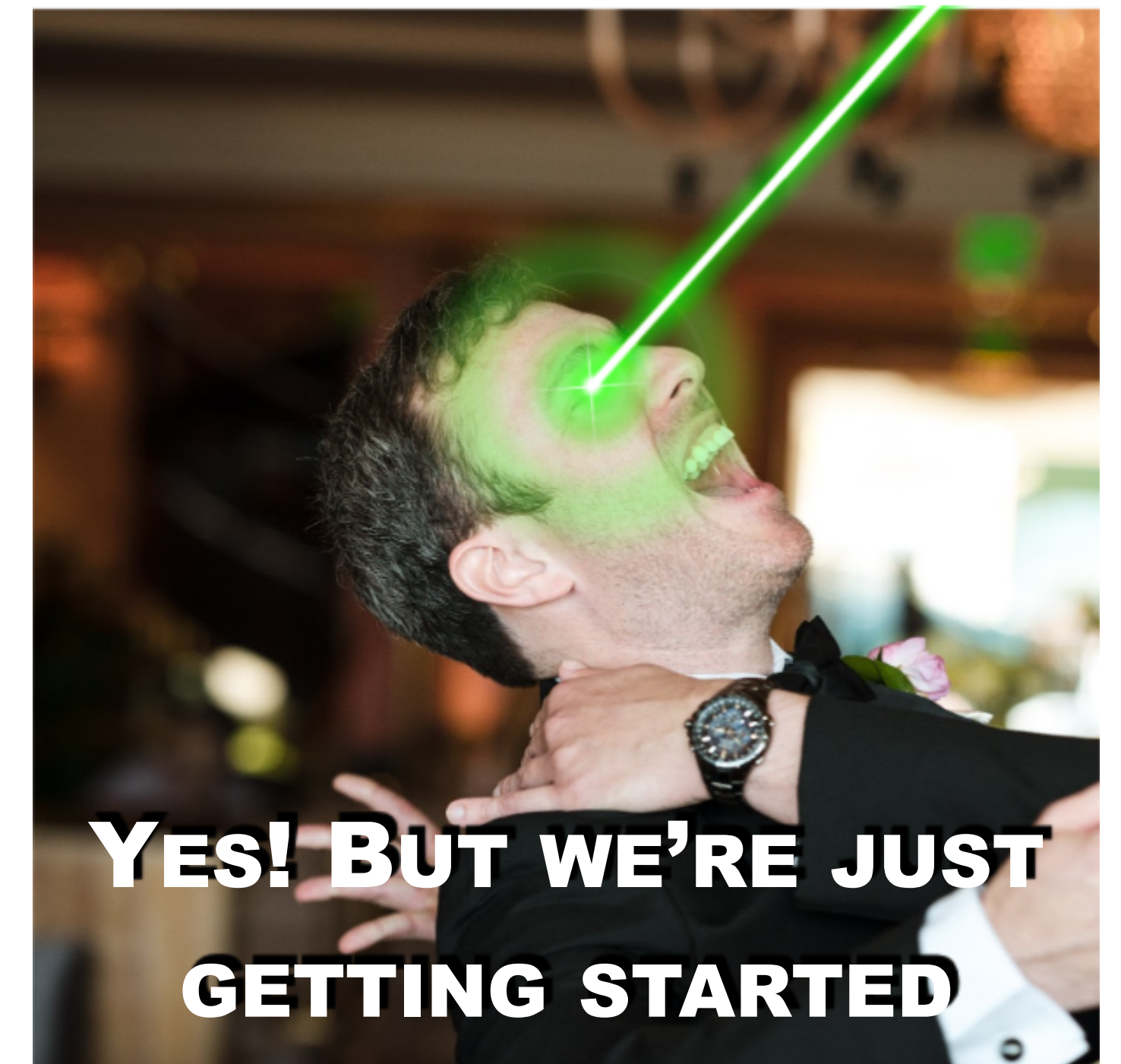
Bretherton, Peters, and Back (2004)

# Precipitation-moisture relation

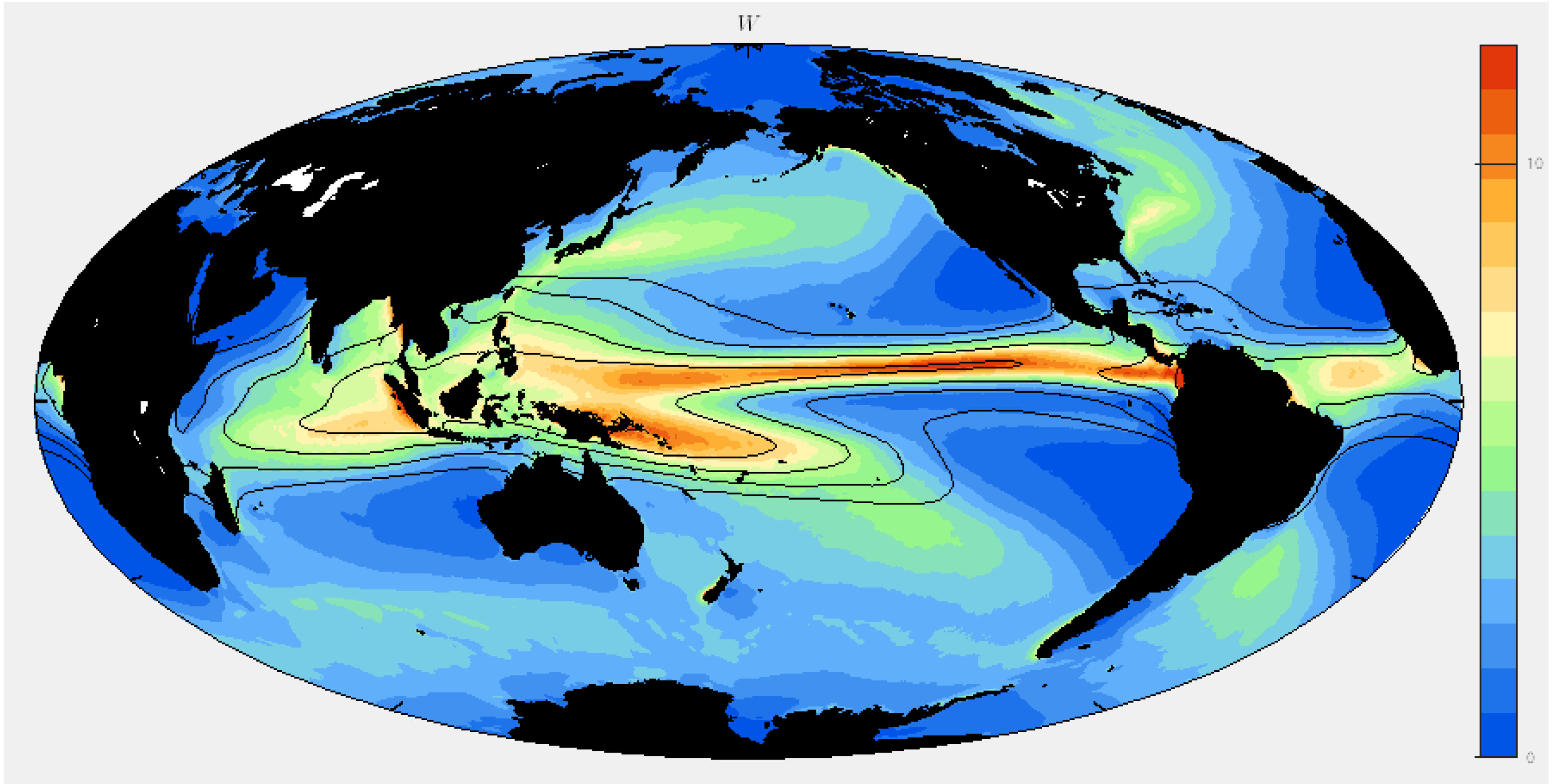


On the surface this seems like an obvious result.

But the fact that it's a Laplacian is interesting. The fact that moisture has a strong control over deep convection under WTG balance will have major implications down the road.

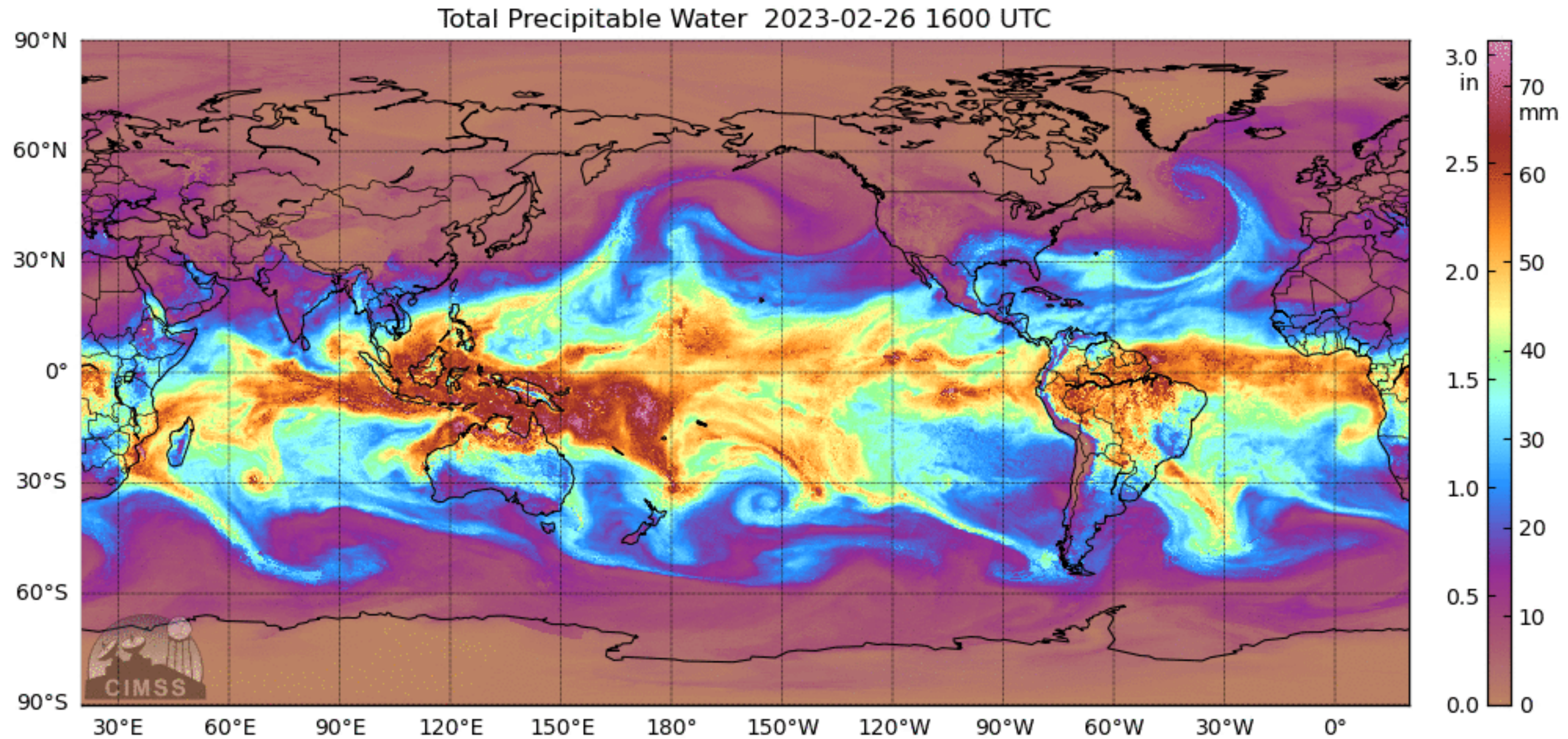


# Precipitation-moisture relation



# Precipitation-moisture relation

A massive consequence is that moving moisture around implies moving rainfall around too.





Draw a schematic of a convective life cycle and point out the main findings of Wolding et al. (2022)

How do the results compare to what you have learned in class?